

M_L^{SM} COMPUTED FROM STRONG MOTION ACCELEROGRAMS RECORDED IN YUGOSLAVIA

V. LEE AND M. TRIFUNAC

Department of Civil Engineering, University of Southern California, Los Angeles, California 90089-2531, U.S.A.

M. HERAK

Geofizički Zavod, Horvatovac b.b. P.O. Box 224, 41000 Zagreb, Yugoslavia

M. ŽIVČIĆ

Seizmološki Zavod Slovenije, Kersnikova 3/II, 61000 Ljubljana, Yugoslavia

AND

D. HERAK

Geofizički Zavod, Horvatovac b.b. P.O. Box 224, 41000 Zagreb, Yugoslavia

SUMMARY

It is shown that the new definition¹ of strong motion local magnitude M_L^{SM} leads to stable estimates of magnitudes for earthquakes in Yugoslavia, with epicentral distances $R < 100$ km and for $2.5 < M_L^{SM} < 6.5$. Tables with magnitudes computed using this new procedure are presented for all earthquakes contributing to the strong motion accelerogram files in EQINFOS for Yugoslavia.² The similarity of our findings with the analogous analyses for California suggests new possibilities for relative calibration between various local magnitude scales, which are used in southeastern Europe, and M_L in California.

INTRODUCTION

In 1935 Richter³ proposed the use of the magnitude scale M_L , which is determined from the peak response of a Wood-Anderson seismometer ($T_n = 0.8$ sec, fraction of critical damping $\zeta = 0.8$ and static magnification $V_s = 2800$). At first this scale was used for epicentral distances greater than 25 km and less than 600 km. Later Gutenberg and Richter⁴ extended its use to small epicentral distances (< 25 km) by using torsional seismometers with static magnification equal to 4. Since it is based only on the peak response, this magnitude cannot characterize any detailed features of the earthquake source. Yet, because it samples amplitudes with frequencies centred near 1 Hz, which is in the centre of the period range of interest to earthquake engineering and strong motion seismology, this scale continues to be useful for many studies of the near source ground motion. With a gradual increase in the number of the recorded accelerograms^{5,6} in California, the associated source mechanism studies^{7,8} and the empirical investigations of the dependence of spectral amplitudes on magnitude,^{9,10} it became possible to investigate the properties of this magnitude scale also at small epicentral distances,¹¹⁻¹⁴ and to show how it can be extended to the estimation of the local magnitude, M_L^{SM} , computed from the recorded strong motion accelerograms.¹

For the estimation of engineering seismic risk the very first step of analysis requires description of the earthquake occurrence in terms of earthquake magnitudes and distances from the site, and then description of the attenuation of strong motion amplitudes with distance. Procedures for detailed computation and mapping of seismic risk are now available for scaling earthquakes both in terms of magnitude and site intensity,¹⁵⁻¹⁷ these having been developed using empirical results on attenuation,¹⁸ local site effects, magnitude and intensity, all in California.^{10,19} For successful transfer of such models elsewhere and for

Table I. Distribution of the available strong motion records versus published magnitude, M_p , and distance $\Delta = (R^2 + H^2)^{1/2}$

M_p	$\log_{10} \Delta$ (km)								
	0.7*	0.9	1.1	1.3	1.5	1.7	1.9	2.1	2.3
6.76-7.25			2	2	1	2	1	9	6
6.26-6.75									
5.76-6.25				3	11	5	3	1	
5.26-5.75		1	3		3	2	3	5	1
4.76-5.25		1	3	9	11	5		1	2
4.26-4.75	4	4	13	28	19	4			1
3.76-4.25	1	3	15	25	12	6	1		
3.26-3.75	1	1	34	10	4	3			
2.76-3.25	2	3	21	6	6	2			

*Intervals 0.2 units wide and centred at 0.7, 0.9, 1.1, . . .

independent development of similar results in other seismically active regions, it is essential first to describe the local magnitude scales and the local attenuation equations in various respective regions.

In the immediate vicinity of the source, the attenuation of the wave amplitudes should be least affected by the quality factor Q , and by other local features of geologic structure, so that M_L^{SM} estimates should be less sensitive to regional geologic and tectonic setting, in contrast to other magnitude estimates using more distant (> 100 km) stations. Therefore M_L^{SM} should also be useful for relative calibration of different magnitude estimates in Europe, for better understanding of the differences caused by various instruments, regional variations of Q and different procedures and magnitude definitions.²⁰⁻²⁴

In 1987 a programme on digitization of 449 accelerograms recorded in Yugoslavia was completed,² contributing digitized data on more than 200 earthquakes between 1975 and 1983. From these data 183 contributing earthquakes have been identified and cross-referenced with various regional catalogues and can be used to compare the regional estimates of magnitude with M_L^{SM} for 325 recorded accelerograms (Tables I and II). The aim of this paper is to present our estimates of M_L^{SM} in Yugoslavia using Trifunac's definition¹ of M_L^{SM} and thus facilitate relative calibration of different more distant estimates of magnitudes in eastern Europe. With those results it will be possible to improve the accuracy of the seismic risk estimates in Yugoslavia and to some extent in the neighbouring Balkan countries.

For small epicentral distances the new definition of M_L^{SM} will take into account the mean effects of the source size, the local geologic conditions and the changes of the shape of the spectral amplitudes of the strong ground motion with magnitude. Since for small epicentral distances most seismological instruments go off scale, the use of M_L^{SM} will provide additional data for estimation of M_L , when this is not available or possible from other sources.

Figure 1 illustrates the data available for this study. It shows that all recordings were obtained between 1975 and 1983, for magnitudes between 2.5 and 7, for mostly shallow focal depths (< 25 km) and for small epicentral distances, typically less than 50 km (Table I). Much of these data came from the Monte Negro earthquake of 1979 and its many aftershocks. The total number of digitized and processed accelerograms is 449. Of those, 325 records could be identified to have occurred during 183 earthquakes.

DEFINITION OF M_L^{SM}

The definition of M_L^{SM} is¹

$$M_L^{SM} = \bar{M}_L^{SM} - D(\bar{M}_L^{SM}) \quad (1)$$

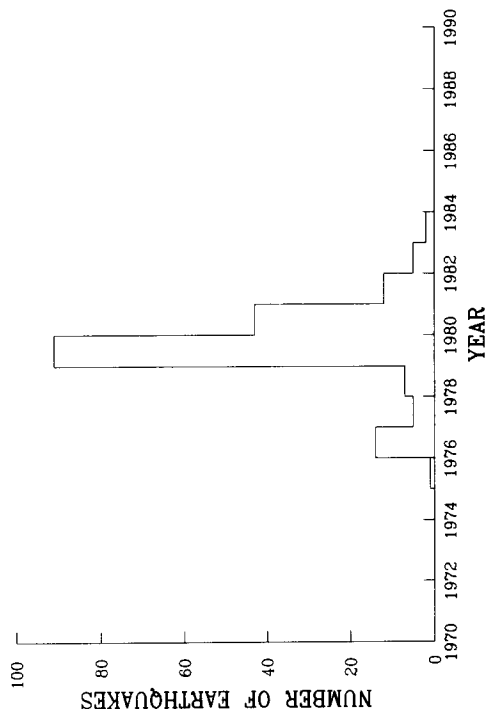


Figure 1(a). Number of earthquakes in the Yugoslav strong motion data base for period from 1975 to 1983

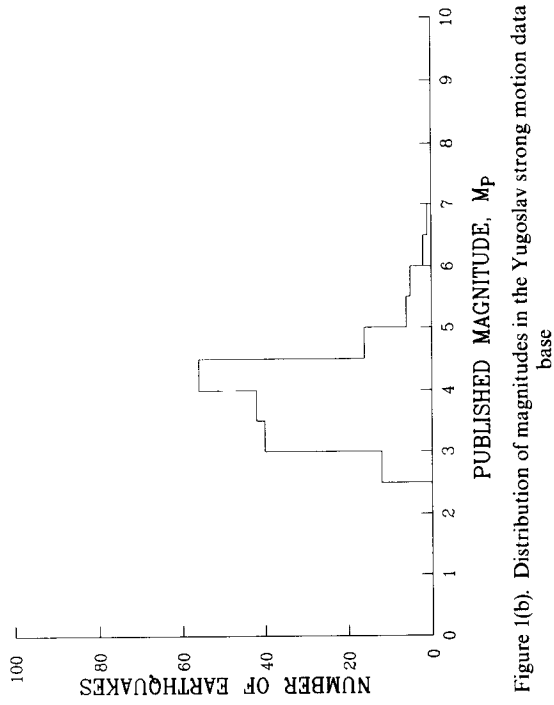


Figure 1(b). Distribution of magnitudes in the Yugoslav strong motion data base

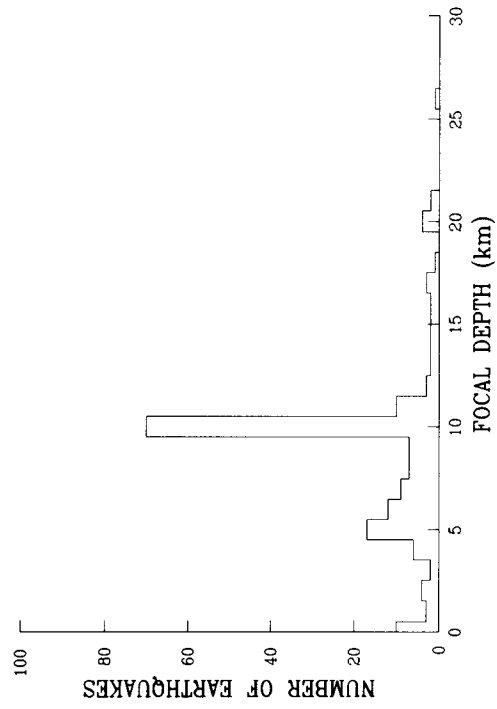


Figure 1(c). Distribution of focal depths

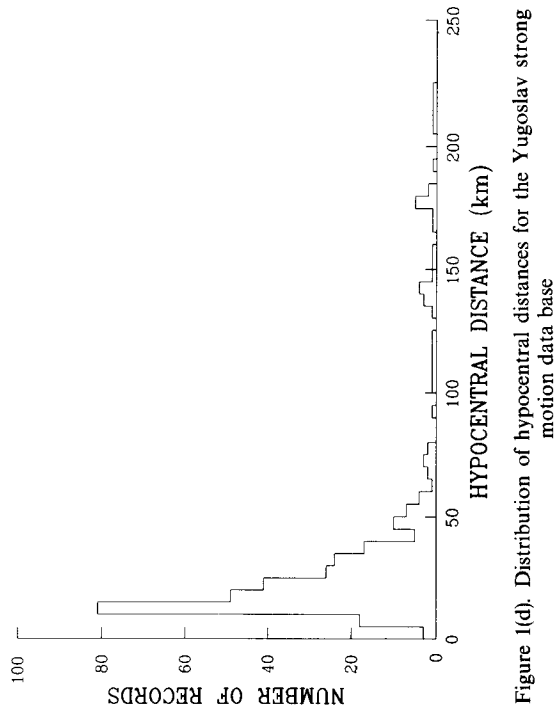


Figure 1(d). Distribution of hypocentral distances for the Yugoslav strong motion data base

Table II.

EQ#	EQ. REF#	DATE		TIME		LAT. (N)	LONG. (E)	H KM	MAGNITUDES					S.M.		# OF	
		MO/DA/YR		GMT	SEC.				1	2	3	4	5	1	2	REC	INT.
1	1	9/ 7/75		1623		45.84	15.74	0	4.50					4.40	4.14		5.5MCS
2	2	5/11/76		0532		46.20	13.11	10	4.10					4.41	4.55	2	5MCS
3	3	2 7/14/76		0539	33.2	46.35	13.31	6	4.40				4.30	3.96	3.47		
4	4	9/ 7/76		1108		46.21	13.02	5	4.10				4.20	3.95	3.68		5.5MCS
5	5	3 9/11/76		1631	11.0	46.30	13.22	7	5.10	4.93		4.93	5.50	4.71	4.13	2	7.5MM
6	6	4 9/11/76		1635	2.8	46.24	13.17	11	5.60				5.70	5.51	5.18	2	9.0MM
7	7	5 9/11/76		1648	54.6	46.29	13.17	1	4.30	3.95		3.95		4.24	4.02		
8	8	6 9/12/76		1953	27.8	46.29	13.24	6	4.50	4.16		4.16	4.50	4.14	3.67		7.0MM
9	9	7 9/13/76		1854	45.9	46.33	13.21	4	4.40	4.33		4.33	4.40	4.45	4.33	2	6.0MM
10	10	8 9/15/76		0315	19.4	46.28	13.18	9	5.80				6.00	5.94	5.82	2	9.0MM
11	11	9 9/15/76		0438	52.7	46.31	13.17	7	4.80	4.57		4.57	4.70	4.55	4.13	2	7.0MM
12	12	10 9/15/76		0458	41.8	46.33	13.21	5	4.60	4.46		4.46	4.40	4.25	3.77		6.0MM
13	13	11 9/15/76		0921	17.8	46.33	13.15	2	6.10				6.10	6.04	5.69	2	9.5MM
14	14	9/15/76		0945		46.30	13.14	5	4.20				4.30	4.24	4.11		
15	15	12 9/15/76		1111	10.8	46.33	13.24	8	5.00	4.67		4.67	4.90	4.65	4.11	2	7.5
16	17	14 9/16/77		2348	7.5	46.28	12.97	11	5.30	4.77		4.77	5.30	4.80	4.09		8.0MM
17	18	15 9/28/77		0143	14.7	46.29	13.05	17	4.20	3.79		3.79		4.32	4.27		
18	19	9/23/77		0258		41.49	20.08	37	4.70					4.75	4.63		7.0MM
19	20	11/ 3/77		0223		42.12	24.03	11	5.50					5.53	5.33		7.MCS
20	21	12/ 7/77		1921		46.26	13.13	10	3.40					3.38	3.34		
21	22	16 1/ 1/78		0423	46.3	43.30	17.60	20	3.80	4.00		4.00		3.52	3.27		5.5MCS
22	23	17 2/20/78		1213	33.5	46.48	13.25	0	4.00	3.62		3.62		4.62	5.08		6.0MCS
23	24	18 3/16/78		0608	39.0	43.10	18.03	20	4.00	3.76	3.98	3.87		3.94	3.76		6.0MCS
24	25	160 6/20/78		2003	25.1	40.75	23.26	20	6.00				6.00	5.95	5.39		
25	26	11/16/78		2023		41.97	21.50	0	3.50					3.59	3.59		5.5MCS
26	28	21 12/17/78		0216	49.3	43.44	17.38	16	4.50					4.22	3.80	2	
27	27	12/31/78		1556		41.99	23.22	21	4.60					5.07	5.37	2	
28	29	22 2/17/79		2206	3.0	44.69	17.25	5	3.40	3.60	3.19	3.40		3.35	3.30		5.5MCS
29	112	23 3/31/79		1555	24.0	41.88	19.07	6	4.00	4.03	4.08	4.06		4.09	4.03	2	
30	113	24 4/ 9/79		0210	19.8	41.90	19.05	0	5.20	5.04		5.04	5.10	5.33	5.26	3	7.5
31	114	4/10/79		1054		41.90	19.19	10	3.60					3.54	3.43		
32	31	26 4/15/79		0619	45.8	42.02	19.07	13	7.00				6.80	7.09	6.74	23	9.0MCS
33	116	27 4/15/79		0631	9.4	42.16	18.74	10	4.90	4.47		4.47		4.87	4.66	3	
34	117	4/15/79		0701		42.00	19.30	10	3.60					3.46	3.32		
35	118	28 4/15/79		0711	27.9	41.98	19.18	11	4.20	3.90	3.74	3.82		4.18	4.00		
36	119	29 4/15/79		0725	31.6	41.94	19.47	10	4.10	4.01	3.66	3.83		4.20	4.14		
37	120	4/15/79		0748		41.80	19.14	10	3.80					3.88	3.82		
38	121	31 4/15/79		0808	41.9	42.24	18.65	10	4.40	4.48	4.32	4.40		4.07	3.64		
39	122	32 4/15/79		0813	17.1	41.92	19.24	26	4.30	4.45	4.35	4.40		4.35	4.23		
40	123	4/15/79		0910		41.92	19.33	10	4.40					4.06	3.62	2	
41	124	37 4/15/79		1025	25.3	1.91	19.40	15	4.90	4.86		4.86		4.85	4.62	3	
42	125	38 4/15/79		1107	30.2	42.08	19.06	6	4.20	3.96	4.10	4.03		4.37	4.37	2	
43	126	39 4/15/79		1142	16.5	41.93	19.41	10	3.70	3.87	3.70	3.79		3.58	3.42		
44	127	40 4/15/79		1243	46.2	42.08	19.19	12	4.40	3.99	4.13	4.06		4.09	3.67	2	
45	128	41 4/15/79		1324	33.6	42.41	18.74	10	4.30		3.97	3.97		3.66	3.13		
46	32	42 4/15/79		1443	5.9	42.26	18.71	9	5.80					5.58	5.11	6	
47	130	4/15/79		1524		42.39	18.87	10	4.50					4.15	3.69		
48	131	43 4/15/79		1752	56.2	42.55	18.55	9	4.10	4.18	3.82	4.00		3.64	3.23		
49	132	46 4/15/79		2049	46.8	42.00	19.21	14	4.30	4.43	4.47	4.45		4.01	3.62		
50	133	49 4/16/79		0756	1.7	41.79	19.56	17	4.20	4.10	4.26	4.18		4.26	4.16		
51	134	51 4/16/79		1004	39.1	41.92	19.24	10	4.90	5.03		5.03		5.17	5.25	2	
52	135	53 4/16/79		1430	51.9	42.00	19.03	15	4.10	4.32	4.09	4.20		4.21	4.17	2	

Table II. (cont.)

EQ#	EQ. REF#	DATE		TIME		LAT. (N)	LONG. (E)	H KM	MAGNITUDES					S.M.		# OF REC	INT.
		MO/DA/YR		GMT	SEC.				1	2	3	4	5	1	2		
53	228	4/16/79		1535		41.82	19.21	14	3.30					4.17	4.85	1	
54	136	55	4/16/79	1551	7.3	41.81	19.37	10	3.80	3.57	3.59	3.58		3.68	3.50	2	
55	137	56	4/16/79	2300	27.0	41.86	19.38	11	4.20	4.30	4.10	4.20		4.48	4.59	3	
56	138	58	4/17/79	0353	32.4	41.80	19.48	10	4.20	3.86	3.95	3.91		4.14	3.93		
57	33	59	4/17/79	0539	57.9	42.45	18.62	0	4.90	5.10		5.10		4.61	4.15	2	
58	140	61	4/17/79	1806	16.0	42.04	19.05	1	4.00	3.92	4.17	4.05		4.22	4.28		
59	141	64	4/18/79	0245	12.1	41.89	19.10	8	3.70	3.96	4.03	3.99		3.63	3.48		
60	142	65	4/18/79	0350	4.9	41.91	19.14	7	4.10	4.17	4.07	4.12		3.95	3.69	3	
61	30	67	4/18/79	1519	20.4	46.34	13.29	21	4.80	4.92		4.92		4.78	4.59	2	7.0MM
62	143	68	4/18/79	1951	13.4	42.01	19.06	5	4.60	4.43		4.43		4.56	4.36	4	
63	144	69	4/19/79	0017	35.3	41.90	19.18	10	4.50	4.42		4.42		4.71	4.75	3	
64	145	70	4/19/79	0542	49.8	42.04	19.03	2	4.60	4.36		4.36		4.63	4.50	4	
65	146	71	4/19/79	0707	5.4	42.01	19.02	3	4.00	4.29	4.15	4.22		3.78	3.51		
66	147		4/20/79	2341		41.88	19.16	10	3.60					3.08	2.81		
67	148		4/21/79	0136		41.80	19.10	5	3.50					3.44	3.36		
68	149		4/21/79	0149		41.85	19.10	5	3.20					3.51	3.67		
69	150	72	4/21/79	0238	5.7	41.98	19.20	20	4.30	4.66		4.66		4.90	5.34	2	
70	151		4/21/79	0404		41.83	19.37	10	3.80					3.73	3.58		
71	152	73	4/21/79	0433	1.2	41.80	19.18	5	4.30	4.19	4.30	4.24		4.17	3.90	2	
72	153	74	4/21/79	0454	26.9	41.83	19.16	5	3.90	3.87	4.27	4.07		3.78	3.57		
73	154		4/22/79	0444		41.95	19.27	10	3.50					2.96	2.72		
74	155	78	4/22/79	0632	12.0	41.92	19.24	5	4.50	4.43	4.45	4.44		4.52	4.39	3	
75	156	79	4/22/79	0732	7.4	41.78	19.32	16	4.00	3.91	3.71	3.81		3.52	3.13		
76	157	80	4/23/79	1252	46.3	41.92	19.26	8	3.20	3.19		3.19		2.81	2.71		
77	158		4/24/79	0023		41.81	19.29	10	3.60					3.53	3.42		
78	159		4/24/79	1645		41.81	19.10	10	3.60					4.10	4.44		
79	160	81	4/24/79	2226	24.9	41.94	19.28	10	3.60	3.54	3.67	3.60		3.37	3.20	2	
80	161	82	4/25/79	0636	46.3	41.92	19.25	6	3.80	3.92	4.08	4.00		3.96	3.96		
81	162	83	4/25/79	1514	32.0	41.83	19.19	11	3.80	3.93	4.07	4.00		3.85	3.77		
82	164		4/25/79	1812		41.94	19.13	5	3.30					3.81	4.13		
83	163	84	4/25/79	1912	17.6	41.85	19.26	11	3.60	3.78	3.92	3.85		3.59	3.51		
84	165	85	4/28/79	0338	2.6	42.19	18.81	2	4.50	4.56		4.56	4.50	4.21	3.79	2	7.0
85	166	86	4/29/79	1024	17.7	42.00	19.23	10	3.90	4.00	4.08	4.04		3.52	3.19		
86	167	87	4/30/79	1700	5.7	42.27	18.82	4	4.50	4.33		4.33		3.72	3.07		
87	168	88	5/ 1/79	0639	6.4	41.93	19.26	4	3.20	3.40	3.41	3.41		3.32	3.38		
88	169	90	5/ 3/79	1639	45.7	41.85	19.13	5	3.50	3.66	3.71	3.68		3.61	3.62		
89	170		5/ 6/79	2252		41.89	19.36	10	3.50					3.23	3.07		
90	34	92	5/12/79	0330	34.2	42.26	18.91	9	5.20	5.08		5.08	5.00	5.25	5.11	6	
91	172	93	5/14/79	0953	7.6	41.93	19.19	10	4.60	4.60		4.60		4.85	4.93	2	
92	173	94	5/20/79	0845	28.5	42.19	18.82	5	4.30	4.59		4.59		4.42	4.38	2	
93	35	95	5/24/79	1723	18.2	42.15	18.76	5	6.20				6.10	6.25	5.99	12	7.5MM
94	175		5/24/79	1942		42.16	18.71	56	3.80					4.15	4.32		
95	176		5/24/79	2228		42.21	18.65	10	4.10					4.00	3.77		
96	177		5/25/79	0332		42.29	18.76	10	3.70					3.93	4.00		
97	178		5/25/79	0722		42.20	18.73	10	4.10					3.65	3.24		
98	179		5/25/79	1145		42.14	18.76	10	4.30					4.23	4.00	2	
99	180		5/27/79	1447		42.15	18.78	10	4.40					4.59	4.59		
100	181		5/28/79	1327		42.12	18.68	10	4.20					4.38	4.39		
101	182		5/30/79	0538		41.85	19.06	10	4.10					4.01	3.79		
102	183		5/30/79	2347		42.30	18.76	10	4.40					3.86	3.33	2	
103	184		6/ 1/79	0929		42.37	18.60	10	3.80					3.68	3.49		
104	185		6/ 4/79	0251		42.13	18.78	8	4.40					4.40	4.24		

Table II. (cont.)

EQ#	EQ. REF#	DATE		TIME		LAT. (N)	LONG. (E)	H KM	MAGNITUDES					S.M.		# OF	
		MO/DA/YR		GMT	SEC.				1	2	3	4	5	1	2	REC	INT.
105	186	6/18/79		0956		42.19	18.65	10	4.30					4.25	4.05		
106	187	6/20/79		2118		42.17	18.69	49	4.80					4.68	4.38	2	
107	188	7/14/79		1407		42.26	18.76	10	3.90					4.27	4.47	2	
108	189	96	7/20/79	0256	2.9	41.86	19.21	5	3.60	3.60	3.53	3.57		4.10	4.43		
109	190		8/ 2/79	1414		42.06	19.04	10	4.20					4.73	5.09		
110	191		8/ 6/79	0748		42.31	18.58	10	4.10					4.12	3.98		
111	192		8/17/79	0530		41.89	19.31	3	4.40					4.30	4.03		
112	193		8/24/79	1016		42.16	18.79	38	3.90					4.04	4.03		
113	36	149	9/ 7/79	1257	56.9	44.84	17.38	0	3.20	3.54	3.15	3.35		2.92	2.85		5.0MCS
114	194	100	9/21/79	1202	41.7	41.95	19.36	4	4.20	3.90	4.17	4.04		4.11	3.86		
115	195		11/ 5/79	1824		42.00	19.31	11	3.20					3.46	3.59		
116	196	101	11/ 6/79	0805	25.6	41.90	19.26	6	3.10		3.65	3.65		3.58	3.87		
117	198	102	11/ 9/79	0148	49.7	41.88	19.30	17	4.10	3.84	4.17	4.00		4.32	4.38		
118	199		11/ 9/79	0238		41.82	19.19	7	3.10					3.49	3.71		
119	200	103	11/ 9/79	0420	2.2	41.87	19.24	4	3.70	3.61	3.57	3.59		3.35	3.10		
120	201	104	11/10/79	0419	34.7	41.90	19.37	6	4.20		4.38	4.38		5.11	5.85		
121	202	105	11/20/79	1831	59.3	42.01	18.96	8	4.50	3.90	3.78	3.84		4.77	4.87	2	
122	37	157	5/18/80	2002	57.9	43.21	20.97	6	5.70					5.87	5.81	8	8.0MCS
123	38	159	5/18/80	2026	42.7	43.24	20.96	11	5.00					5.07	4.96	2	
124	39		5/18/80	2019		43.26	20.90	10	4.30					4.56	4.66		
125	40		5/18/80	2041		43.29	20.89	1	4.90					4.81	4.53		
126	42	106	5/21/80	0922	41.0	43.33	21.00	7	3.80	3.90	3.79	3.85		3.52	3.26		6.0MM
127	43	107	5/23/80	1226	23.9	43.28	21.04	0	4.50	4.58		4.58		4.81	4.96		
128	44	108	5/23/80	1237	35.5	43.19	21.02	5	3.20	3.28		3.28		3.48	3.62		
129	45		5/23/80	1340		43.12	21.30	10	3.00					3.44	3.71		
130	46	109	5/25/80	0603	36.2	43.26	21.06	13	3.60	3.72	3.79	3.76		4.14	4.51		
131	47	110	5/25/80	0708	49.4	43.27	20.95	9	3.60	3.58	3.22	3.40		3.53	3.41		
132	48	111	5/26/80	0025	37.1	42.87	20.98	10	3.40	3.07		3.07		3.32	3.26		
133	50		5/31/80	1642		43.30	20.80	10	2.80					2.98	3.17		
134	51	113	6/ 1/80	2124	44.6	43.28	21.01	8	3.40	3.53	3.50	3.51		3.93	4.27		
135	52	114	6/ 3/80	1908	5.7	43.25	21.00	10	3.50	3.50	3.34	3.42		3.58	3.57		
136	53	115	6/ 4/80	0321	43.3	43.27	20.99	4	3.10	2.94		2.94		3.30	3.43		
137	54		6/ 4/80	2129		43.31	20.80	10	3.20					3.21	3.23		
138	55		6/ 5/80	0603		43.27	21.00	10	3.00					3.13	3.26		
139	56	116	6/ 9/80	0811	22.7	43.07	20.73	10	2.90					3.21	3.43		
140	57	117	6/10/80	2125	1.8	43.34	21.05	10	4.10	4.02	3.95	3.99		4.07	3.89	2	
141	58	118	6/12/80	2346	26.7	43.05	20.98	12	3.50	3.49	3.05	3.27		4.11	4.54		
142	59		6/14/80	0642		43.01	20.61	10	3.30					3.97	4.43		
143	60	119	6/14/80	0220	21.6	43.05	20.97	18	3.30	3.07		3.07		3.53	3.62		
144	61	120	6/17/80	0952	6.0	43.25	20.95	10	3.70	3.80	3.57	3.68		3.96	4.04	2	
145	62	121	6/17/80	2214	39.5	43.23	20.95	10	3.70	3.84	3.50	3.67		3.81	3.80	2	
146	63	122	6/19/80	0147	2.3	43.19	20.92	10	3.60	3.57	3.55	3.56		3.95	4.13		
147	64	123	6/19/80	0442	16.1	43.27	21.07	10	3.50	3.31	3.47	3.39		3.28	3.14		
148	65	124	6/28/80	0610	12.0	43.22	20.94	6	3.40	3.27	3.16	3.21		3.59	3.66		
149	66	125	6/29/80	0552	11.7	43.21	20.88	9	3.70	3.84	3.49	3.66		3.49	3.30		
150	68	126	7/ 1/80	0643	11.4	43.26	20.96	9	3.10		2.99	2.99		3.03	3.05		
151	69	127	7/ 2/80	1423	52.1	43.27	20.97	8	3.10					2.91	3.02		
152	70		7/13/80	2054		43.29	20.63	10	3.10					3.82	4.33		
153	71	128	7/13/80	2207	53.3	44.73	17.31	7	3.40		2.96	2.96		2.95	2.77	4	
154	41	129	7/19/80	0037	57.9	41.45	20.38	12	4.50	4.80		4.80		4.72	4.86		?6.5MCS
155	72	130	7/31/80	2152	39.3	43.31	20.95	10	3.20	3.44	3.42	3.43		3.61	3.84		

Table II. (cont.)

EQ#	EQ. REF#	DATE		TIME		LAT. LONG.		H	MAGNITUDES					S.M.		# OF	
		MO/DA/YR		GMT	SEC.	(N)	(E)		1	2	3	4	5	1	2	REC	INT.
156	74 132	9/ 3/80		1159	40.9	43.26	20.93	10	3.20	3.51	3.26	3.39		3.29	3.34		
157	75 134	10/10/80		0103	46.7	43.23	20.91	10	2.80					2.85	3.00		
158	76 135	10/11/80		1055	12.0	43.28	20.95	10	3.00					2.80	2.81		
159	77	10/11/80		2339		43.20	20.30	0	2.90					3.28	3.52		
160	78 136	10/21/80		1943	11.1	43.23	20.88	10	3.90					3.93	3.82	2	
161	79 151	11/ 3/80		1911	45.7	43.22	20.85	10	3.80		3.68	3.68		3.68	3.49	2	
162	81 152	12/ 8/80		0632	0.9	43.31	21.03	10	3.30	3.51		3.51		3.41	3.45		
163	82 153	12/14/80		0254	48.6	43.27	21.07	10	3.90	4.10	4.28	4.19		4.22	4.38	2	
164	83 154	12/22/80		1909	39.9	43.22	20.99	10	3.70	3.88	3.49	3.68		3.93	4.00	2	
165	87	2/28/81		2253		42.95	20.56	0	3.90					4.00	3.96	2	
166	91 156	3/ 7/81		0653	16.0	42.95	20.78	0	3.80	3.74	3.82	3.78		3.70	3.52		
167	92	3/ 8/81		1310		42.84	20.68	10	3.40					3.54	3.56		
168	98	5/11/81		1325		43.27	18.53	10	3.40					3.53	3.55		5.5MCS
169	99 137	7/24/81		0253	43.1	44.71	17.27	5	3.00	3.15	3.14	3.14		2.70	2.67	3	
170	100 138	7/24/81		0255	51.6	44.67	17.24	10	2.90	2.90	2.79	2.84		2.67	2.71	2	
171	101 139	8/13/81		0258	13.5	44.70	17.22	7	5.40					5.89	6.16	4	8.0MCS
172	102 140	8/13/81		0437	12.4	44.69	17.19	7	3.50	3.56		3.56		2.97	2.73	2	
173	103 142	8/14/81		0444	54.4	44.73	17.22	10	3.20	3.32	3.19	3.29		2.95	2.89	2	
174	204	8/19/81		2043		42.17	18.95	10	4.50					4.39	4.13		
175	104	8/21/81		0330		44.89	17.37	11	3.20					3.30	3.36	3	
176	105	8/30/81		0311		44.98	17.40	10	2.80					2.82	2.97	2	
177	110	6/ 2/82		0542		43.35	20.94	2	4.6					5.04	5.30		7.5MCS
178	106 144	7/ 3/82		0341	32.5	44.68	17.19	6	2.80		3.18	3.18		2.66	2.75	3	5.0MCS
179	109 158	7/14/82		1614	53.8	42.13	21.43	7	4.40		4.42	4.42		3.97	3.48	3	6.0MCS
180	107 146	10/12/82		0133	59.3	44.69	17.14	6	3.30					3.20	3.16	4	5.0?MM
181	108	11/22/82		1857		44.58	16.80	10	2.9					3.49	3.88	3	
182	203 147	1/ 5/83		0403	30.5	41.96	19.19	10	3.80		3.98	3.98		4.38	4.78		
183	111 148	2/25/83		1822	13.6	41.95	21.66	6	4.50		4.17	4.17		4.29	3.93	9	6.5MCS

* REPORTED FOCAL DEPTHS ARE NOT RELIABLE

where

$$\bar{M}_L^{SM} = M_L^* - b_2(M_L^*)(2 - s) \quad (2)$$

and

$$M_L^* = \log_{10} A_{\text{synthetic}} - \text{Att}(\Delta_0) \quad (3)$$

$A_{\text{synthetic}}$ is the peak amplitude (in mm) of the computed response of a Wood-Anderson seismometer, the attenuation function $\text{Att}(\Delta_0)$ is given¹ in Table III for $\Delta_0 = (R^2 + H^2)^{1/2}$, and R and H are the epicentral distance and the source depth respectively. $D(\bar{M}_L^{SM})$ is given in Tables IV(a) and IV(b) and represents the mean deviation of \bar{M}_L^{SM} from M_L or M_p (median of local magnitudes published in various seismological catalogues) computed from distant stations, but for earthquakes recorded by strong motion accelerographs in Yugoslavia (Table II). When M_p or M_L is available one should use $D(\bar{M}_L^{SM})$ defined versus M_p as in Table IV(a). When M_L or M_p is not available $D(\bar{M}_L^{SM})$ is evaluated from Table IV(b) where it is given in terms of \bar{M}_L^{SM} . $b_2(M_L^*)$ represents a correction function for average local site amplification effects. $s = 0$ corresponds to a site on sediments, while $s = 2$ is for sites on geological basement rocks.¹⁶ The amplitudes of $b_2(M_L^*)$ are given in Table IV(c). The above computation of M_L^{SM} is carried out for each recorded horizontal acceleration component. All such estimates are then averaged to give M_L^{SM} for the earthquake.

Table III. Att (Δ_0) for $\Delta_0 = \sqrt{R^2 + H^2}$ *

Δ_0 (km)	Att (Δ_0)	Δ_0 (km)	Att (Δ_0)	Δ_0 (km)	Att (Δ_0)
1.0	-1.62	100.0	-3.08	340.0	-4.21
5.0	-2.08	120.0	-3.13	360.0	-4.30
10.0	-2.30	130.0	-3.18	380.0	-4.38
15.0	-2.42	140.0	-3.23	400.0	-4.45
20.0	-2.51	150.0	-3.28	420.0	-4.52
25.0	-2.58	160.0	-3.33	440.0	-4.58
30.0	-2.63	170.0	-3.38	460.0	-4.63
35.0	-2.68	180.0	-3.43	480.0	-4.69
40.0	-2.71	190.0	-3.48	500.0	-4.73
45.0	-2.75	200.0	-3.53	520.0	-4.78
50.0	-2.78	220.0	-3.63	540.0	-4.82
60.0	-2.83	240.0	-3.73	560.0	-4.85
70.0	-2.88	260.0	-3.83	580.0	-4.88
80.0	-2.93	280.0	-3.93	600.0	-4.90
90.0	-2.98	300.0	-4.02		
100.0	-3.03	320.0	-4.12		

*This attenuation law has been proposed on the basis of recordings made mostly in California.¹

Table IV(a). $D(\bar{M}_L^{\text{SM}})$ in equation (1) versus M_p , for data in Yugoslavia

M_p	2.5	3.0	3.5	4.0	4.5	5.0	5.5	6.0	6.5	7.0
$D(\bar{M}_L^{\text{SM}})$	1.91	1.69	1.45	1.21	0.96	0.70	0.43	0.15	-0.14	-0.43

Table IV(b). $D(\bar{M}_L^{\text{SM}})$ in equation (1) versus \bar{M}_L^{SM} , for data in Yugoslavia

\bar{M}_L^{SM}	4.0	4.5	5.0	5.5	6.0	6.25	6.50	6.75	7.00	7.25
$D(\bar{M}_L^{\text{SM}})$	1.84	1.67	1.48	1.08	0.58	0.33	0.08	-0.17	-0.42	-0.67

Table IV(c). $b_2(M_L^*)$ in equation (2), for data in California

M_L^*	3.5	4.5	5.5	6.5	7.5
$b_2(M_L^*)$	0.10	0.10	0.11	0.12	0.13

It is seen that equation (3) is analogous to

$$M_L = \log_{10} A - \log_{10} A_0(R) \quad (4)$$

which corresponds to the original definition of the local magnitude scale,^{3,25} with A (in mm) representing the peak response of a Wood-Anderson seismometer, and $\log_{10} A_0(R)$ representing Richter's empirical attenuation law.

M_L^{SM} FOR DATA IN EQINFOS FILES

By using the new definition of the local magnitude scale M_L , as described above, we have computed M_L^{SM} for all earthquakes which contributed to the uniformly processed strong motion data base in EQINFOS files for Yugoslavia.² Table II presents the list of the contributing earthquakes.

Using equations (2) and (3) we evaluated \bar{M}_L^{SM} for each record associated with an earthquake in Table II. Empirical relations showing the frequency dependent attenuation on sediments and on basement rock^{10,18} are not available for Yugoslavia. In this analysis we assume that $b_2(M_L^*)$ in equation (2), as determined for California data,¹ can be applied here as well. Site characterization for Yugoslav data is also not available. Since most recordings there have been obtained on sediments, in this analysis we will adopt an approximate working assumption that $s = 0$ for all data.¹⁶ Then by comparing the estimates of \bar{M}_L^{SM} with the published magnitudes for these events we derived $D(\bar{M}_L^{SM})$, which is shown in Figure 2. $D(\bar{M}_L^{SM})$ was evaluated by using the Richter^{3,25} attenuation law given by $\log_{10} A_0(R)$ and also by using the new attenuation function in the near field 'Att (Δ_0)' proposed by Trifunac,¹ 'without constraints' and 'with constraints' (that the source dimension S and the source coherence length S_0 must be less than the hypocentral distance Δ). The data on $D(\bar{M}_L^{SM})$ were smoothed by averaging over 0.5 magnitude intervals and plotted as points in Figure 2. Tables IV(a) and IV(b) represent the curve 'with constraints' for data in Yugoslavia (Figure 2) and show results versus M_p and versus \bar{M}_L^{SM} . M_L^{SM} can be evaluated either using $D(\bar{M}_L^{SM})$ versus M_p or versus \bar{M}_L^{SM} , but the estimates using M_p are more stable because of saturation of M_L^{SM} for large events. Figure 3(a) shows M_L^{SM} versus M_p using $D(\bar{M}_L^{SM})$ in terms of M_p and for all 183 available events, containing many one point estimates. The resulting standard deviation of the estimates is equal to 0.30 magnitude units. Considering only those estimates for which 2 and more than 2 strong motion records are available, the number of M_L^{SM} estimates reduces to 67, but the standard deviation also reduces to 0.26 magnitude units. Figure 3(b) shows M_L^{SM} versus M_p estimates using $D(\bar{M}_L^{SM})$ in terms of \bar{M}_L^{SM} . The standard deviation of these estimates is 0.55 for 183 data points. Again, ignoring those M_L^{SM} estimates that are based on only one strong motion recording would again reduce the number of data points to 67 and the standard deviation of the estimates to 0.50. It is seen that the estimates in Figure 3(a) are better.

Column 1 in Table II shows the earthquake number (1–183) of the 183 earthquakes used in the analysis. Column 2 is the cross-reference with the earthquake number listed in Table I of EQINFOS files.² Column 3 gives the corresponding cross-reference earthquake number listed in Table A2 of the Yugoslav EQINFOS data.² Column 4 shows the date of the earthquake given in the order MONTH/DAY/YEAR. Column 5 presents the time of the earthquake in GMT to the nearest minute. Columns 6 and 7 show the latitude and

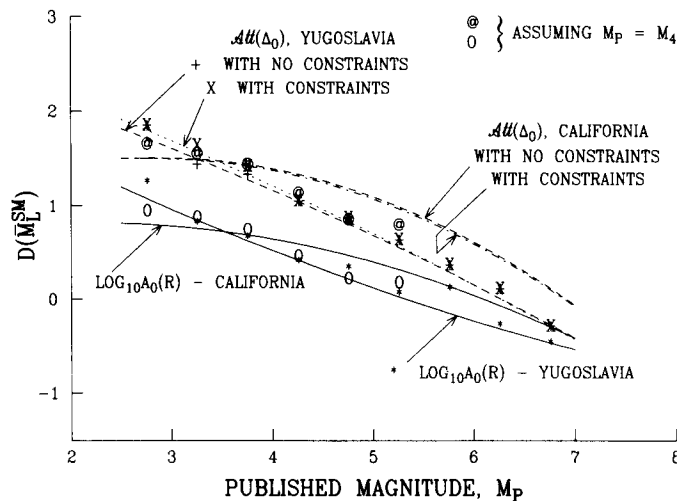
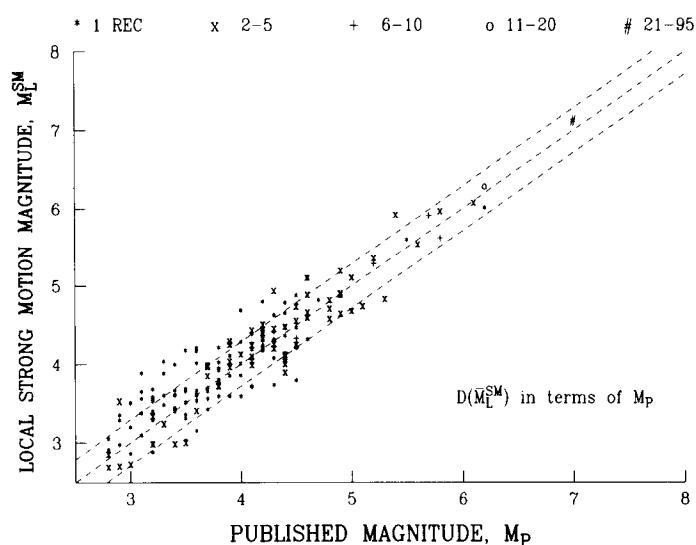
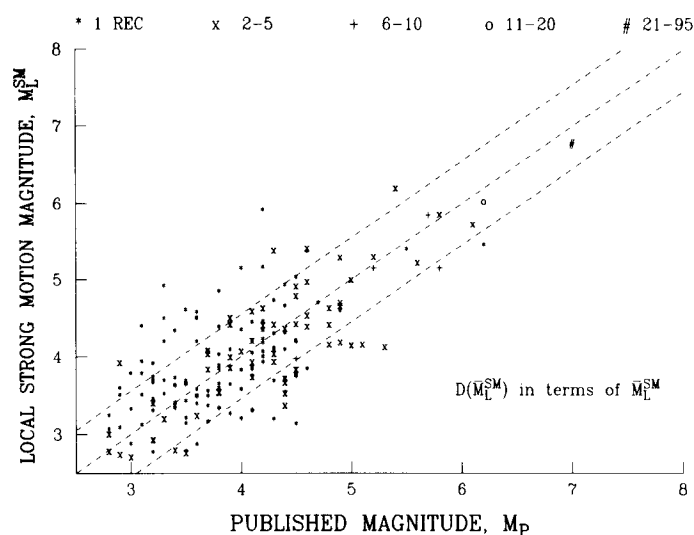


Figure 2. Comparison of $D(\bar{M}_L^{SM})$ for strong motion data in Yugoslavia and California, versus published magnitude, M_p

Figure 3(a). M_L^{SM} versus M_p , using $D(\bar{M}_L^{SM})$ in terms of M_p Figure 3(b). M_L^{SM} versus M_p , using $D(\bar{M}_L^{SM})$ in terms of \bar{M}_L^{SM}

the longitude of the epicentre. Column 8 gives the focal depth in km. All events with entries in column 3 of Table II were located using all the available data including the local networks (see Appendix A in Jordanovski *et al.*²). The coordinates of the hypocentres for the rest of the events were taken from ISC and CSEM bulletins. Column 9 lists the published (1) magnitude M_p . Columns 10 to 13 present different magnitude estimates (2,3,4 and 5) as discussed in the following text. Columns 14 and 15 give the strong motion magnitudes M_L^{SM} calculated using equations (1) to (3) (with $D(\bar{M}_L^{SM})$ versus $M_p(1)$ and $M_L^{SM}(2)$, respectively). Column 16 shows the number of strong motion stations used in the computation of M_L^{SM} (only for those events with two or more stations recording the strong motion). Column 17 gives the epicentral intensity.

The magnitudes in column 9 of Table II represent median values of local magnitudes published by ISC and CSEM. To test how homogeneous that data may be we carried out the following study. The data recorded at

two Yugoslav stations VIR (on the island Vir near Zadar) and LJU (Ljubljana) between 1975 and 1983 were analysed. VIR was in operation from 1975 to 1982 and recorded most events from Table II. Both VIR and LJU were equipped with short period seismographs ($T_0 = 0.5$ to 1.0 sec) with record amplitudes proportional to the ground velocity. Since our analysis showed that the period of the peak response did not vary much, the records at these two stations were scaled with respect to the reference set of events given in Jordanovski *et al.*² (Table II), using

$$M = \log_{10} A_{\max} + p \log_{10} R + q$$

where M is local magnitude, A_{\max} is maximum record amplitude in mm, and R is epicentral distance in km. At VIR half peak to peak amplitudes of P_g and S_g waves, $A(P_g)$ and $A(S_g)$, were read together with their corresponding periods $T(P_g)$ and $T(S_g)$. Using the above equation, different magnitude estimates M_{i^*} were then determined after replacing A_{\max} by: $A(S_g)$ for $i^* = 1$, $A(P_g)$ for $i^* = 2$, $[A(S_g)A(P_g)]^{1/2}$ for $i^* = 3$, $A(S_g)/T(S_g)$ for $i^* = 4$ and $A(P_g)/T(P_g)$ for $i^* = 5$, and regressing for p and q . At LJU the peak amplitude A_{\max} and its period T_{\max} were read and two magnitudes M_{j^*} were determined, using A_{\max} for $j^* = 1$ and A_{\max}/T_{\max} for $j^* = 2$.

For completeness of this presentation, Table II also shows all data of Table A2 in Jordanovski *et al.*² The third column of this table presents the earthquake number used in Table A2. The tenth and eleventh columns give M_2 (averages of M_{i^*} , $i^* = 1, \dots, 5$) and M_3 (average of M_{j^*} , $j^* = 1, 2$) respectively. The twelfth column gives M_4 which is the average of M_2 and M_3 , when both are available, or equals M_2 or M_3 when either of the two is available. Thus, by comparing M_1 and M_4 , one can study the consistency of almost all the magnitude estimates in our data base. Out of 160 events present in Table A2 of Jordanovski *et al.*² 137 can be compared in this way.

The thirteenth column in Table II presents M_5 , which corresponds to the estimates of M_L contained in Table I of Ambraseys.²² It is seen that his data base contains 19 events in Friuli, Greece and Yugoslavia, for the period considered by our study. Those are typically events with $M_L \geq 4.2$, and their magnitudes agree favourably with our estimates designated by M_1 , in the ninth column of Table II. Typical differences are about 0.1 magnitude units.

To test the stability of the M_L^{SM} estimates we computed $D(\bar{M}_L^{SM})$ again for all the data associated with these 137 events, now assuming that M_4 is equal to M_p . The results are shown as points in Figure 2. It is seen that the estimates based on M_4 lead to essentially the same results as the large data set in Table II, using M_p as median of M_L from different European stations. We conclude that $D(\bar{M}_L^{SM})$ as shown in Figure 2 is sufficiently stable and can be used for computation of M_L^{SM} in Yugoslavia.

DISCUSSION AND CONCLUSIONS

By computing \bar{M}_L^{SM} from equation (2) and evaluating $D(\bar{M}_L^{SM})$ for strong motion accelerations recorded in Yugoslavia we found that the final estimates of M_L^{SM} in this region are as stable as the corresponding estimates in California (the standard deviation of the estimates of \bar{M}_L^{SM} in Figure 3(a) is 0.3 magnitude units for all the 183 earthquakes in Table II). Estimates using $D(\bar{M}_L^{SM})$ determined from Table IV(a), in terms of M_p , give more stable results on M_L^{SM} than estimates using Table IV(b), with $D(\bar{M}_L^{SM})$ in terms of \bar{M}_L^{SM} .

We employed standard and published algorithms for epicentral and focal determinations and in most cases with an adequate number of stations and data points. Typical error of such locations (for events identified in column 3 of Table II and originally listed in Table A2 of Jordanovski *et al.*²) is about 5 km. Then, the dependence of M_L^{SM} on hypocentral distance is contained in the attenuation function (Table III), indicating that, for distances greater than ~ 10 km, uncertainties in focal locations do not contribute much to the uncertainties in the computed magnitudes (less than 0.1 magnitude units).

We studied $M_L^{SM} - M_p$ versus distances R and $\Delta = (R^2 + H^2)^{1/2}$ (where R is the epicentral distance and H is the focal depth) and found no systematic departures from zero (see also Luco,¹³ Trifunac¹). This implies that the new attenuation function $Att(\Delta_0)$ can be used for Yugoslav earthquakes without any modification

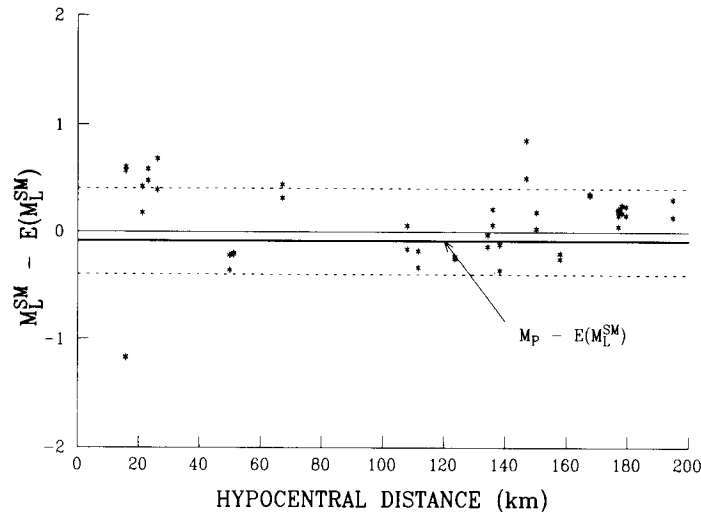


Figure 4. $M_L^{SM} - E(M_L^{SM})$ versus hypocentral distance

being indicated by the currently available data for $\Delta < 100$ km. For the Monte Negro earthquake of April 15, 1979 ($M_p = 7$ and $M_L^{SM} = 7.1$) $M_L^{SM} - E(M_L^{SM})$ estimates for the individual stations (E designates expected value) were symmetrically scattered about zero for Δ up to about 200 km, suggesting that our extension¹ of $Att(\Delta_0)$ beyond 100 km, using the shape^{3,25} of $\log_{10} A_0(R)$, is also in agreement with the recorded data (Figure 4).

Trifunac¹ considered several causes and possible explanations for the amplitudes and the shape of $D(\bar{M}_L^{SM})$ versus M_p . At present it appears that for M_p less than about 5.5 to 6, the amplitudes of $D(\bar{M}_L^{SM})$ are governed by the Q structure in the region, as sampled by the strong motion data and by more distant records which are used to estimate M_p . The similarity of the shape of $D(\bar{M}_L^{SM})$ in California¹ and in Yugoslavia and no obvious dependence of $M_L^{SM} - M_p$ versus Δ for data in Yugoslavia suggest that the nature of the attenuation of the strong motion amplitudes in these two different regions of the world is similar. Trifunac and Todorovska²⁶ arrived at similar conclusions using independent comparisons of the intensity data. This means that it will be possible to use California data in seismic risk studies in Yugoslavia (or vice versa) more directly and without major corrections for magnitude scaling and near field attenuation. $Att(\Delta_0)$ represents attenuation in a narrow frequency band centred around periods equal to 1 sec. Therefore, for example, for use of frequency dependent scaling relationships on spectral amplitudes¹⁰ developed for California data, still further analyses and verification will be required before those can be used in Yugoslavia.

Figure 2 also shows the amplitudes of $D(\bar{M}_L^{SM})$ for California data. It is seen that for small magnitudes, near 3.5, $D(\bar{M}_L^{SM})$ for California data is nearly the same as $D(\bar{M}_L^{SM})$ estimated for Yugoslavia. However, between magnitudes 4.5 and 6.5 California data give $D(\bar{M}_L^{SM})$ larger than $D(\bar{M}_L^{SM})$ for Yugoslavia by 0.4 to 0.5 magnitude units. This suggests a somewhat larger overall average value of Q for Yugoslavia (say $Q \sim 150-200$) than for California ($Q \sim 100-150$) assuming that (i) the average distances between strong motion and distant recordings are comparable and also (ii) that the distant recordings can be assumed to lead to the same estimates as the corresponding stations would in California. Detailed analysis and interpretation of such trends is beyond the scope of this paper, as this requires large scale analysis of many recordings in Europe. In this work, as in other studies for California,¹ we want to determine only the relative scaling of the recorded strong motion amplitudes and of magnitudes determined by the regional seismological networks. This will enable correct relative scaling of recorded strong motion amplitudes, and then the data in various catalogues, which are used in seismic risk studies and mapping, can be used with greater confidence and with more accuracy.

REFERENCES

1. M. D. Trifunac, ' M_L^{SM} ', *Soil dyn. earthquake eng.* (in press) (1990).
2. L. Jordanovski, V. W. Lee, T. Olumčeva, C. Sinadinovski, M. Todorovska and M. D. Trifunac, 'Strong earthquake ground motion data in EQINFOS, Yugoslavia, Part I', Department of Civil Engineering, *Report No. 87-05*, University of Southern California, Los Angeles, CA, 1987.
3. C. F. Richter, 'An instrumental earthquake scale', *Bull. seism. soc. Am.* **25**, 1–32 (1935).
4. B. Gutenberg, and C. F. Richter, 'Earthquake magnitude, intensity, energy and acceleration', *Bull. seism. soc. Am.* **32**, 163–191 (1942).
5. V. W. Lee, and M. D. Trifunac, 'EQINFOS (The strong motion earthquake data information system)', Department of Civil Engineering, *Report No. 82-01*, University of Southern California, Los Angeles, CA, 1982.
6. M. D. Trifunac and V. W. Lee, 'Uniformly processed strong earthquake ground accelerations in the Western United States of America for the period from 1933 to 1971: Corrected acceleration, velocity and displacement curves', Department of Civil Engineering, *Report No. 78-01*, University of Southern California, Los Angeles, CA, 1978.
7. M. D. Trifunac, 'A three-dimensional dislocation model for the San Fernando, California, earthquake of February 9, 1971', *Bull. seism. soc. Am.* **64**, 149–172 (1974).
8. M. D. Trifunac, 'Inversion of earthquake source mechanism using near field strong motion data', *Int. conf. comput. methods exper. measurements*, Capri, Italy, 65–75 (1989).
9. M. D. Trifunac, 'Preliminary empirical model for scaling Fourier amplitude spectra of strong motion acceleration in terms of earthquake magnitude, source to station distance and recording site condition', *Bull. seism. soc. Am.* **66**, 1343–1373 (1976).
10. M. D. Trifunac, 'Dependence of Fourier spectrum amplitudes of recorded strong earthquake accelerations on magnitude, local soil conditions, and on depth of sediments', *Earthquake eng. struct. dyn.* **18**, 999–1016 (1989).
11. B. A. Bolt, 'The local magnitude M_L of the Kern County earthquake of July 21, 1952', *Bull. seism. soc. Am.* **68**, 513–515 (1978).
12. H. Kanamori and P. C. Jennings, 'Determination of local magnitude, M_L , from strong-motion accelerograms', *Bull. seism. soc. Am.* **68**, 471–485 (1978).
13. J. E. Luco, 'A note on near-source estimates of local magnitude', *Bull. seism. soc. Am.* **72**, 941–958 (1982).
14. M. D. Trifunac and A. G. Brady, 'On correlation of seismoscope response with earthquake magnitude and modified Mercalli intensity', *Bull. seism. soc. Am.* **65**, 307–321 (1975).
15. M. D. Trifunac, 'Seismic microzonation mapping via uniform risk spectra', *Proc. 9th world conf. earthquake eng.* Tokyo-Kyoto, Japan **VIII**, 78–80 (1988).
16. M. D. Trifunac and A. G. Brady, 'On the correlation of seismic intensity scales with the peaks of recorded strong ground motion', *Bull. seism. soc. Am.* **65**, 139–162 (1975).
17. M. D. Trifunac and V. W. Lee, 'Uniform risk spectra of strong earthquake ground motion', Department of Civil Engineering *Report No. 85-05*, University of Southern California, Los Angeles, CA, 1985.
18. M. D. Trifunac and V. W. Lee, 'Frequency dependent attenuation of strong earthquake ground motion', *Soil dyn. earthquake eng.* **9**, 3–15 (1990).
19. M. D. Trifunac, 'Preliminary empirical model for scaling Fourier amplitude spectra of strong motion acceleration in terms of modified Mercalli intensity and geologic site conditions', *Earthquake eng. struct. dyn.* **7**, 63–74 (1979).
20. N. N. Ambraseys, 'Magnitude assesment of Northwestern European earthquakes', *Earthquake eng. struct. dyn.* **13**, 307–320 (1985).
21. N. N. Ambraseys, 'The seismicity of Western Scandinavia', *Earthquake eng. struct. dyn.* **13**, 361–399 (1985).
22. N. N. Ambraseys, 'Uniform magnitude re-evaluation of European earthquakes associated with strong motion records', *Earthquake eng. struct. dyn.* **19**, 1–20 (1990).
23. A. A. Kiratzi and B. C. Papazachos, 'Magnitude scales for earthquakes in Greece', *Bull. seism. soc. Am.* **74**, 969–985 (1984).
24. J. Vanek, A. Zatopek, V. Kárník, N. Kondorskaya, Y. Riznichenko, E. Savarenskij, S. Solovov and N. Shebalin, 'Standardization of magnitude scales', *Izv. akad. nauk SSSR ser. geofiz.* **2**, 153–158 (1962).
25. C. F. Richter, *Elementary Seismology*, Freeman, San Francisco, 1958.
26. M. D. Trifunac, and M. I. Todorovska, 'Attenuation of seismic intensity in Albania and Yugoslavia', *Earthquake eng. struct. dyn.* **18**, 617–631 (1989).